



## PETROLOGY OF GABBROIDES AND ISOTOPE SIGNATURE OF SULFIDE MINERALIZATION FROM FEDOROV-PANSKY LAYERED MAFIC INTRUSION, KOLA PENINSULA, RUSSIA

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**Abstract:** Samples from Fiodoro-Pansky massif were investigated by polarized light microscopy, EDS electron microscopy and isotope ratio mass spectrometry. We have observed four-stage ore mineralization: (1) magmatic Fe-Ti-oxides, (2) sulfide-ferrous oxide, (3) hydrothermal Cu-Ni-sulfide with haycockite, talnakhite and galena admixture, (4) tiosulfates formed from pyrrhotite. A few sulfide samples from generations (2) and (4) were investigated isotopically. The second generation is isotopically homogenous with  $\delta^{34}\text{S}$  close to zero, whilst the third generation shows significantly negative  $\delta^{34}\text{S}$  values.

**Keywords:** Fedorov-Pansky massif, layered intrusion, Kola, petrology, ore mineralization, sulfides, stable isotopes.

### 1. INTRODUCTION

In the middle part of Kola Peninsula, northern part of Baltic shield, is located the Imandra-Varzuga belt. This belt is constituted of metamorphised sedimentary-volcanic rock complex.

Towards North from this belt, Kola rock formation of metamorphised granite-gneiss is located, see **Fig. 1**. Between these two large formations the Fiodorov-Pansky (F-P) massif is located. The massif consists of layered mafic intrusions of gabbroide rocks. The age of this intrusions ranges from 2501 to 2446 Ma (Latypov and Chistyakova, 2000; Schissel *et al.*, 2002).

The F-P massif belong to the province of the following mafic intrusions: Tornio, Näränkävaa, Penikat, Koillismaa, Lukkulaivaara which are also distributed in Sweden, Finland and Karelia (Iljina and Hansky, 2005; Sharkov, 2006). Within the whole province, mafic layered intrusions are of similar age and structure. Platinum

group elements (PGE) often occur in the rocks of these intrusions. The average of PGE content in the F-P massif is 0.1-0.5 ppm, with the Pd/Pt ratio from 4.2 to 3.3 (Latypov and Chistyakova, 2000).

In this study we analyzed rock samples collected from outcrops and cores of boreholes located in the F-P massif.

### 2. METHODS

Rock samples were analyzed on a HITACHI S-4700 with FESEM (Field Emission Scanning Electron Microscopy) electron microscope equipped with energy dispersion spectroscope (EDS) following by petrography investigation in transmitted and reflected light. Selected mineral samples of chalcopyrite and pyrite were analyzed by means of mass spectrometry for sulfur isotope composition.

The EDS analyses were performed on Field Emission Electron Scanning Microscopy in the Institute of Geology at the Jagiellonian University, Krakow. Petrographic analyses were conducted in the Institute of Earth Sciences and in the Faculty of Geology, Geophysics and Environmental Protection at the Mining Academy (AGH) in

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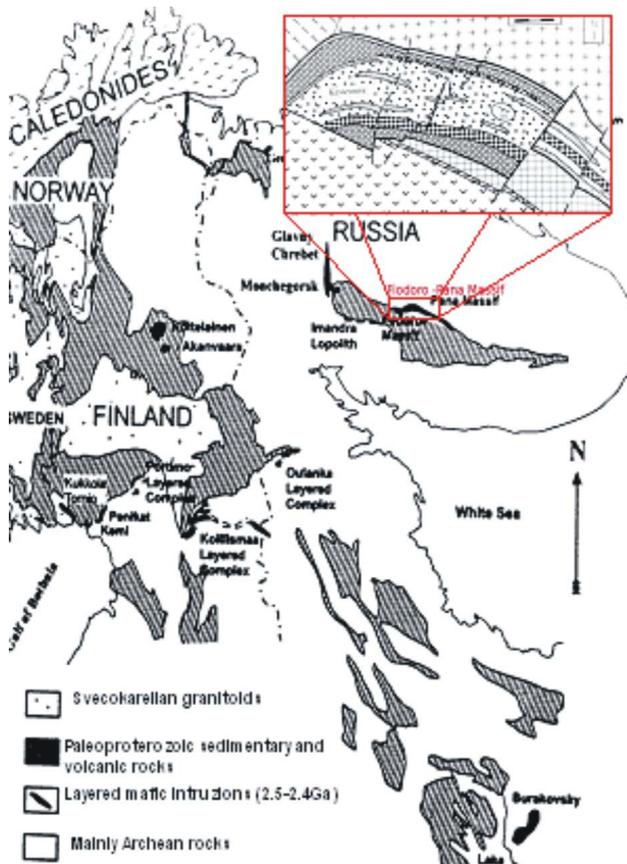


Fig. 1. Geological map showing position of intrusive rocks (after Mitrofanov et al. 2005, Latypov and Chistyakova, 2000).

Krakow, whilst the isotope analyses were performed in the Institute of Physics at Marie Curie-Skłodowska University (UMCS), Lublin. The isotope analysis was done on dual inlet and triple collector mass spectrometer using  $\text{SO}_2$  prepared off line by CuO oxidation. The precision of  $\delta^{34}\text{S}$  values was 0.05‰, the results were normalized to the VCDT scale by analysis  $\text{SO}_2$  gas prepared from the international standard IAEA-S-1.

### 3. RESULTS AND DISCUSSION

The F-P massif mainly consists of layered gabbro-norites, with contain olivine-norite and olivine plagiopyroxene horizons with Cu, Ni -sulfide mineralisation (Mitrofanov et al., 2005, Iljina and Hansky, 2005). The F-P massif consists of several zones (Shissel et al., 2005), namely: marginal mafic schists (bottom zone), taxitic gabbro-norite (bpCa), norite (bCpa, pbC), main gabbro-norite (paCb), lower layered horizon (bC, pC, pbaC), gabbro zone (paCb, pbaC), upper layered horizon (po-Cab), upper gabbro zone (paCb).

These rocks are composed of partly sericitized plagioclase, mainly bytovnite (Table 1). Pyroxenes occur between plagioclase crystals forming offite type of structure. In some samples, oligoclase and garnet of grossular composition (83.93% grs, 16.07%alm) have been detected. Within pyrrhotitic type of gabbro xenolites of ultramafic rocks, e.g. norite composed of cpx, opx, olivine (Fig. 2b), have also been described (Latypov and

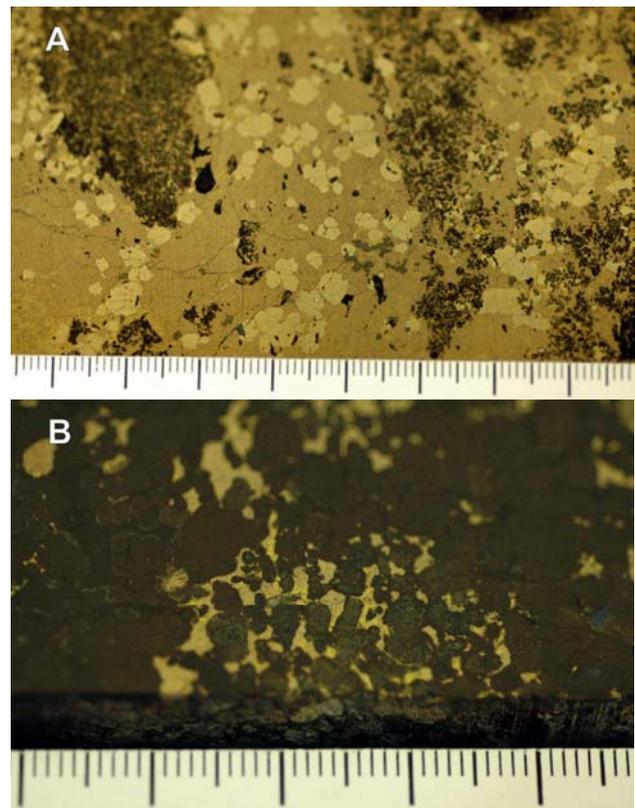


Fig. 2. (a) Macrophotographs of massive sulfate ores (pyrite + pyrrhotite), (b) gabbro containing some pyrrhotite.

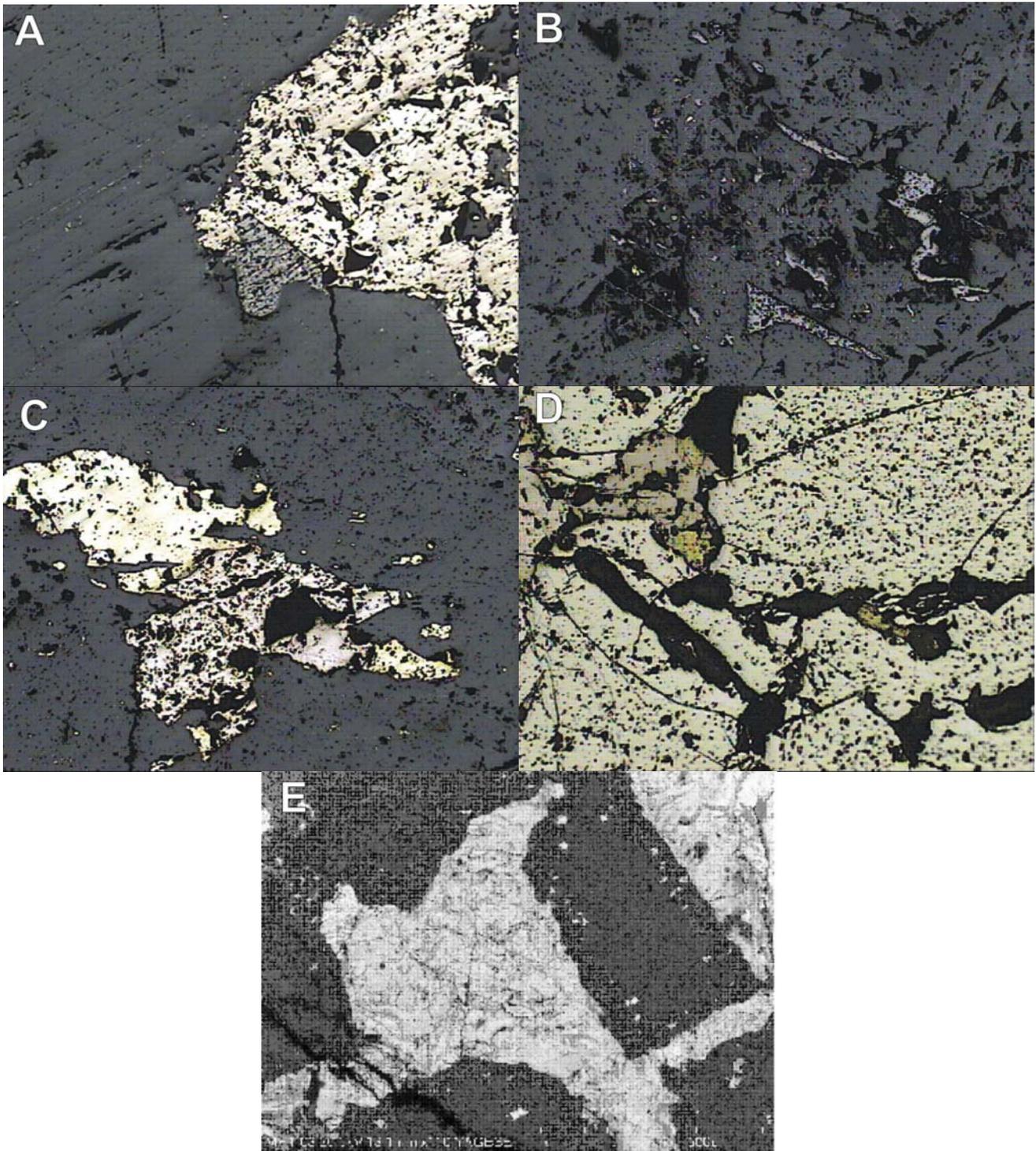
Chistyakova, 2000). Massive sulfate ore shoots composed of pyrite, chalcopyrite and pentlandite occurred in the layer complex (Fig. 2a).

Ore mineralization was developed during multistage processes. The first stage is related to gabbroidal type of host rocks. This stage is characterized by presence of magnetite, titanomagnetite and ilmenite (Fig. 3a). The next stage is characterized by sulphide association, that seems to be younger in comparison to oxide one, because sulfides cut oxides. Sulfides are represented by pentlandite, pyrrhotite, pyrite, chalcopyrite, haycockite, talnakhite and minor galena. Sulfides formed dispersions, veinlets and locally massive nest type structures. Usually formed intergrowth between themselves. Chalcopyrite group of sulfides is younger than Fe-Ni- association. Pyrrhotite is locally oxidized to thiosulfates. (Fig. 3e).

The isotopic composition of the investigated sulfides expressed in terms of delta values is shown in Table 2. It should be noted that, with one exception, the results are close to zero in the VCDT scale, i.e. the isotopic composition is like in meteorites and primary Earth crust formed before 2.2 Ga (Hattori et al., 1983). The only exception is the sample 490.183\_8-1 with more negative value. However, this sulfide was formed later than the remain-

Table 1. Chemical composition of plagioclase based on EDS measurements.

Sample	1MP	1MP	4MP	5MP
% An	82.1	89.16	80.44	77.75



**Fig. 3.** Microphotographs (in reflected light, 1N) of ore minerals: (a) pentlandite with magnetite, (b) vainlets structure of magnetite, (c) pentlandite and chalcopyrite, (d) chalcopyrite, pyrite and pyrrhotite in massive ore, (e) sulfides with oxidized to thiosulfates zone (BSE microphotographs).

ing sulfides, most likely prior to 2.2 Ga ago, when sulfur was already fractionated isotopically by oxidative-reductive reactions (Monster *et al.*, 1983, Hattori *et al.*, 1983).

A multistage character of ore mineralization is a result of genesis of the host rocks. The oldest generation composed of Ti-Fe-oxides is related to the stage of a cumulate formation. Probably pentlandite can be included to end of this stage. The next stage of deposit formation, composed of Cu-Fe-Ni sulfides, is probably related to the hydro-

thermal activity during interaction of a primitive magma injected to the chamber with the magma which underwent differentiation. The same source of sulfur is confirmed by the similar sulfur isotope composition. Small differences in  $\delta^{34}\text{S}$  may result from differentiation during fluid transport. Some sulfides from hydrothermal association are partly oxidized. As a result of such a process, thiosulfate phases have been formed. This process is developed mostly on pyrrhotite.

**Table 2.**  $\delta^{34}\text{S}$  isotopes of chalcopyrite and pyrite.

Sample	05MP	06MP	490.183_8-1	350/64_2	55759/5690a
$\delta^{34}\text{S}_{\text{VCDT}}$ (‰)	0.16	0.28	-1.98	0.60	-0.30

The sulfur isotopic composition of the investigated sulfides points on their origin from the magma injected before 2.2 Ga and also thereafter (sample 490.183\_8-1). Hence, we may conclude from our isotope data that the magmatic process in the F-P massif was multistage. In contrast to our results, the sulfur isotopic composition published by Schissel *et al.* (2002) are significantly scattered, most likely due to lower precision of isotope analysis (0.3‰). Both the deposit and host rocks strata have been affected by regional type of metamorphism of the green facies, which is confirmed by presence of albite-chlorite association and, grossularite.

#### 4. CONCLUSION

Among the four mineralogical stages on the basis of the stable isotopic analysis we have demonstrated the two-stage formation of sulfide mineralization in gabbroide rocks in F-P massif. The first stage (mineralogical stage 2) is magmatic, whilst the second (mineralogical stage 3) is hydrothermal. The older sulfide generation is isotopically correlated with the age of host rock.

#### ACKNOWLEDGEMENTS

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